

Research Article

Hydrological regime and forest development have indirect effects on soil fauna feeding activity in Central European hardwood floodplain forests

Nicole Scheunemann¹, David J. Russell¹

1 Senckenberg Museum of Natural History Görlitz, Department of Soil Zoology, Am Museum 1, 02826 Görlitz, Germany Corresponding author: Nicole Scheunemann (nicole.scheunemann@senckenberg.de)



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Abstract

Soil fauna act as regulators of decomposition processes via their feeding activity, thereby playing an important role in regulating carbon cycling and sequestration. Hardwood floodplain forests are critically endangered habitats, but strongly contribute to carbon sequestration in Central Europe. In the present study, within a floodplain forest-development programme, we investigated the feeding activity of soil fauna via the Bait Lamina test in hardwood floodplain forests of the middle Elbe River in Germany in sites with different hydrological regimes and forest-development stages, with neighbouring grassland sites for comparison.

While statistically significant differences in overall feeding activity between general hydrological regimes or forest development stages were not found, decreases in feeding activity with soil depth were strongly modulated by these factors, indicating more unfavourable conditions for soil fauna at increasing soil depth due to, e.g., anoxic conditions in floodplains of tributaries or low soil moisture content below the shallow rooting zone of grasslands. Registered effects of soil texture on soil fauna feeding activity were dependent on forest-tree density, and combined effects indicate that soil-fauna feeding activity varies with soil temperature during spring, but with soil moisture in early autumn.

In conclusion, our results highlight the importance of the current abiotic conditions on soil-fauna feeding activities in floodplain forests, i.e. soil temperature, moisture and ground water level. Hydrological regime and forest development have a strong impact on the effect of these conditions, indirectly affecting soil fauna feeding activity and highlighting the multifactorial influence on soil fauna functional activity to be considered in floodplain-forest restoration programs.

Key words: Bait Lamina, drivers, soil, soil depth, soil moisture, soil texture

Introduction

Central European hardwood floodplain forests are endangered ecosystems that are characterized by high biodiversity (Tockner et al. 2002) and carbon sequestration ability (Giese et al. 2003; Heger et al. 2021). Inundation events occurring more or less regularly shape these ecosystems and their species communities (Ward et al. 1999; Russell and Griegel 2006; Glaeser and Wulf

2009). However, the area covered by hardwood floodplain forests has declined considerably over the last centuries due to, on the one hand, river management for floodwater protection (e.g., river regulation, dyke building) and, on the other hand, deforestation and agricultural use (Koenzen 2009, 2021).

Along the Middle Elbe River of Germany, efforts are being undertaken to restore natural hardwood floodplain forests, primarily by re-establishing natural oak forests, among other measures (Damm 2013). Thereby, effects of forest restoration on biodiversity and, especially, important ecosystem services are being evaluated. Important factors influencing restoration activities are the hydrological river regime of specific restoration sites as well as different forms of forest re-establishment (hereafter called forest development stages). The influence of soil type and other soil parameters on forest establishment and the associated impacts on biodiversity and ecosystem services are of special importance in these studies.

Floodplains are hotspots of carbon sequestration (Walling et al. 2006; Kiss et al. 2015; Shupe et al. 2022), an important ecosystem service in light of global climate change. Thereby, soils are the basis of important carbon-cycle processes such as degradation of organic matter, mineralisation to inorganic compounds as well as carbon sequestration (e.g., Comerford et al. 2013; Adhikari and Hartemink 2016; Saccá et al. 2017). Soil mesofauna (i.e. soil fauna of 0.2–2 mm body size such as Collembola and Enchytraeids) generally act as regulators of these processes, i.e. by influencing the activity of soil microorganisms via their feeding activity (Bardgett 2005; Crowther et al. 2011; Filser et al. 2016). A high feeding activity of soil fauna is usually accompanied by heightened activity of microorganisms (i.e., growth, enzyme production, nutrient translocation; Tordoff et al. 2011) and positive effects on carbon cycling (Geissen and Brümmer 1999; Römbke 2014; Zhang et al. 2015).

The Bait Lamina test is a well-established and frequently used method for estimating feeding activity of soil mesofauna (Sulkava et al. 1996; Helling et al. 1998; Gongalsky et al. 2008) by providing food bait, the loss of which being an indicator of faunal feeding activity (von Törne 1990; Kratz 1998). The method provides insight into the activity patterns of soil fauna along a depth gradient covering the uppermost 8 centimetres of soil (von Törne 1990; Römbke 2014). Soil fauna activity is not necessarily connected to soil fauna densities or species richness, as even large numbers of, e.g., Collembola might be rather inactive depending on the temporal habitat conditions (Gongalsky et al. 2004). Therefore, mesofauna densities are not a proxy for feeding activities, but rather feeding activity of the existing fauna appears to be regulated to a large extent by prevailing abiotic factors. Several field studies investigated the differences in soil fauna feeding activity between seasons, some suggesting that high temperatures result in highest activities in summer in spite of lower soil moisture (Gongalsky et al. 2008; Rozen et al. 2010; Krishna and Mohan 2017), while others highlight the importance of soil moisture as the main driver of soil-fauna feeding activity (Simpson et al. 2012). However, to our knowledge, the effects of other soil properties on soil fauna feeding activity are widely unknown (Geissen et al. 2007; van Gestel et al. 2009; Birkhofer et al. 2011; Klimek et al. 2015), although their effect on the soil ecosystem itself might be tremendous (Bardgett 2005; Frouz 2018).

In the present study, we investigated soil fauna feeding activity in German hardwood floodplain forests along the middle Elbe River. Corresponding to the studied factors in the forest-restoration programme, we compared forest sites with different hydrological regime, i.e. in the active Elbe floodplain affected by sporadic inundation events, in the seepage water zone behind the Elbe dykes that were historically flooded but currently are only affected by high groundwater levels, and nearby tributaries usually independent from the main Elbe hydrological regime. Further, within the active floodplain, we assessed differences in the soil-faunal feeding activities between forests of different age and tree density (reflecting forest-restoration methods) and compared them to neighbouring agriculturally managed grasslands as the starting point of renewed forest restoration (referred to as forest development stages in the following). We also considered various soil properties, vegetation and soil fauna in the analyses to assess potential drivers of soil fauna feeding activity in the floodplain forests and their potential effects within forest development stages or hydrological regimes.

We hypothesized that soil fauna feeding activity in the studied hardwood floodplain forests is affected (1) by hydrological regime with lower activity rates in potentially water saturated active floodplain soils as compared to drained soils behind the dykes, and (2) by forest development stage with activity rates increasing from grasslands to young and further to old forests. In addition, we hypothesize that (3) differing environmental conditions within these gradients will influence the overall gradient effects on soil fauna feeding activities.

Methods

Study site and sampling design

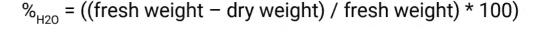
The study area is part of the UNESCO Middle Elbe Biosphere Reserve in northern Germany, and covers ca. 100 km of the Elbe River. The area is characterized by a Central European temperate climate with a mean annual temperature of 9.3 °C and mean annual precipitation of 615 mm (measured at the Lenzen weather station, 53.08°N, 11.48°E). The Elbe floodplain is an anthropogenically altered landscape with a history of dyking, deforestation and agricultural use, with the studied hardwood floodplain forests representing small remnants of the historically extended floodplain forest ecosystem. The arboral vegetation in the study sites is typical for central European hardwood floodplain forests and is characterized by oak (*Quercus robur*) and elm (*Ulmus laevis*); typical understory vegetation is hawthorn (*Crataegus monogyna*). For further information on the study area see Shupe et al. (2021).

In the research programme in which the current study took place, two gradients of ecological habitats are being investigated: a gradient of hydrological regimes and a forest development gradient. For the hydrological regimes, nine forest sites in the study area were investigated (Fig. 1), of which three were located in the active floodplain (i.e. the Elbe floodplain area in front of a dyke flooded during sporadic inundation events, usually after snowmelt in early spring; although none took place during the study period), three in the seepage water zone (i.e. the zone behind the Elbe dykes currently affected by occasional high groundwater levels), and three in the floodplains of tributaries independent from the active Elbe flooding regime. For the forest development gradient, a total of 16 sites of the re-forestation programme were investigated, all of which

are located in the active floodplain (Fig. 1). Of these sites four represented dense older forest stands (mean tree age 80–200 years with multi-layered canopy cover and a rich shrub layer), four were in sparse older forest stands (mean tree age 80–200 years with a less developed canopy and shrub layer), and four established in young forest plantations (mean tree age 17–26 years with a strong grass-dominated herbaceous layer). Additionally, four sites were located in nearby agricultural grasslands and served as a control, since these represented the starting point of reforestation. Old dense, old sparse, young forests and grassland are referred to as forest development stages in the following.

Environmental data on vegetation (e.g., cover and species number of vegetation layers, litter cover and thickness, leaf area index) were collected by botanical project partners (see Suppl. material 1. Mean tree age was estimated by measuring annual tree rings of four dominant oak trees per forest site (Shupe et al. 2021). Data on soil properties (e.g., pH, texture, C and N content) of individual forest sites, as well as precipitation data were obtained from soil-science project partners (Vásconez Navas et al. 2023, see Suppl. material 1). The average number of flooding days per year as a 25-year-mean (1990–2016) was estimated for the individual active floodplain forest sites using a digital model developed for the Elbe river by Weber and Hatz (2020) and Weber and Rosenzweig (2020). Density and species number of soil fauna as used as explanatory variables were obtained from 10 MacFadyen soil cores per forest site (Scheunemann et al., in prep.).

Soil temperature and moisture were assessed at the starting day of each Bait Lamina test at 10 measurement points per forest site. Soil temperature at 10 cm depth was measured using a digital soil thermometer, while soil moisture was calculated from soil cores up to 5 cm depth taken adjacent to the soil thermometer. These soil cores had been used for extraction of soil fauna by heat (data not presented, Scheunemann et al., in prep.). The fresh soil cores were weighed, then dried during animal extraction over 10 days at maximally 55 °C and weighed again. The gravimetric water content was calculated as follows:



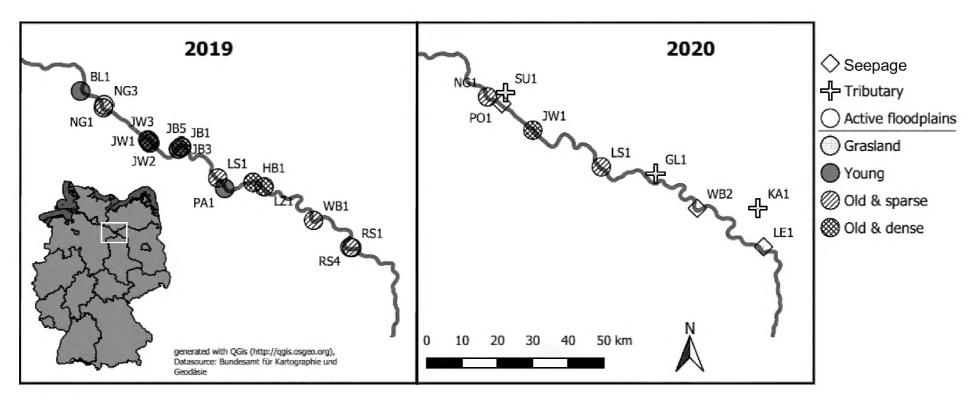


Figure 1. Map of sampling sites along the Elbe river investigated in 2019 (forest development gradient, left) and 2020 (hydrological regimes, right).

Calibration with additional drying at 105 °C for two days revealed essentially no difference to the method above. Mean values of soil temperature and moisture measurements per forest site and season were then used as explanatory variables in the statistical analyses.

Bait lamina test

Feeding activity of soil mesofauna was assessed by the Bait Lamina test according to the ISO standardisation 18311:2016 (ISO 18311 2018). Test strips $(12 \times 0.5 \times 0.1 \text{ cm})$ are made of PVC and contain a row of 16 drilled holes of 1 mm diameter at a distance of 5 mm from each other. The holes were filled with a mixture of powdered cellulose (70%), wheat bran (27%; < 500 μ m particle size) and activated charcoal (3%) as standard bait for soil invertebrates. Filled Bait Lamina strips were obtained from terra protecta GmbH, Berlin, Germany.

We measured soil fauna feeding activity along the forest development gradient in 2019 and in the hydrological regimes in 2020 (see below for detailed sampling dates). The Bait Lamina tests were carried out at the forest sites in three plots at each site during the average periods of highest soil-faunal activity (and avoiding summer dry periods) in late spring and early autumn to assess seasonal differences. Distance between plots within a forest site was a minimum of 10 m to avoid autocorrelation (Irmler 1998), distance between the two closest sites was minimally 300 m (between forest and comparative grassland sites) and maximally 10 km (between active floodplain and tributary sites; Fig. 1). Test strips were placed vertically in the soil such that the bait holes were located at soil depths between 1 and 8 cm to allow evaluation of bait-material disappearance in the respective soil depths. In each plot of every site, 16 test strips were placed in the soil in a 4 × 4 grid with 10 cm distance between single strips. Test strips were exposed to the soil for four weeks. In the forest-development gradient, the experiment started in spring between April 29 and May 3 2019 and finished on May 27/28 2019, and again started in early autumn between September 23 and 26 2019 and finished on October 21/22 2019. In the hydrological regimes, the experiment started in spring on May 11/12 2020 and finished on June 8/9 2020, while in early autumn it started on September 22/23 2020 and finished on October 20/21 2020.

After removal from the field, test strips of each plot were collected together, placed in a separate plastic bag and frozen at -20 °C to stop microbial decomposition of the bait material until further processing. In the lab, attached soil was carefully removed from defrosted strips and the strips visually inspected for empty bait holes by holding against a diapositive slide viewer as a light source. As stipulated by ISO 18311:2016, data was collected as "actively fed on" if light could be observed to shine through a bait hole, and as "not fed on" with bait holes through which no light emitted (even if small feeding traces were found at the sides). Every individual hole in each test strip was inspected in this way, resulting in a data set consisting of binary data representing the feeding activity at each Bait Lamina position (= soil depth). Strips that had been disturbed (chewed on or removed from soil by wild animals) during the four-week field exposure were excluded from further analyses. From all 16 bait laminae of one plot the percentage of bait holes that had been fed on was calculated for every soil depth. Further, the results of the three plots per forest site were averaged to avoid pseudoreplication.

Statistical analyses

A number of numerical environmental parameters (i.e. soil properties, vegetation and soil fauna parameters; see Suppl. material 1) were available for all forest sites. In a first model-building step, these were analysed for collinearity, which was performed separately for both datasets (forest development gradient 2019 and hydrological regime 2020), since these partially consisted of different forest sites. Collinearity of numerical parameters was investigated by principal component analysis using the prcomp() function in the "vegan" package (Oksanen et al. 2020) after standardization to zero mean and unit variance using the decostand() function. To achieve dimension reduction, we used the resulting site scores of the first three principle components (PC; all that explained more than 10% of variance) instead of the original 42 site parameters as independent variables in the subsequent analyses. Inspection of the scores of individual environmental factors showed that, in the 2020 data set (hydrological regime), PC1 represented the hydrological regime (i.e., number of flooding days per year), PC2 represented soil parameters and PC3 represented vegetation parameters as well as soil pH (Suppl. material 2). In contrast, for the 2019 data set (forest development), the first PC mainly reflected vegetation parameters (e.g. species numbers of shrubs and trees, leaf area index, litter cover), PC2 reflected soil parameters (e.g., total C and N content, pH, soil moisture) and PC3 reflected soil fauna density and further vegetation parameters (Suppl. material 3). Further, we built a correlation matrix using the corrplot() function in the "corrplot" package in R studio (Wei and Simko 2021, not presented). The correlation matrices were inspected for correlation indices, and factors with correlation indices > 0.5 or <-0.5 were considered to be autocorrelated. We calculated correlation indices using Pearson's as well as Spearman's correlation coefficient, which both showed similar results. As expected after the PCA, we found correlations within vegetation parameters, soil properties and soil fauna parameters, supporting the reduction of the measured environmental data to the respective PCs.

As categorical variables we used hydrological regime (active floodplain, seepage water zone, tributary) or forest development (old dense, old sparse, young forest, grassland), as well as soil texture estimated a priori by visual inspection in the field (the categories sandy and loamy corresponded well with sand content measured in the lab). In case we found high correlation of a categorical variable with the respective PC (GVIF value > 5 after using the vif() function of the "car" package; Fox and Weisberg 2019), we excluded the respective PC from further analyses and only retained the categorical variable. In this way, in the hydrological regime data set PC1 was replaced by the categorical "hydrology" variable, and in the forest development data set the categorical variable "forest development" replaced PC1.

We then applied linear mixed effects models (Imer() in the "Ime4" package (Bates et al. 2015)) to the bait-lamina data sets of the mean percentage of "fed upon" baits at a respective soil depth as dependent variable and site identity and season as random terms. The null models only included soil depth and the two random terms as independent variables, while hydrology (or forest development), soil texture, the respective PCs as well as all possible two-fold interactions of these parameters were added consecutively to the model (forward selection) by hand. Based on Akaike's Information Criterion (AIC) and anova() functions in the "stats" package (implemented in R), after every step models were compared and only the better model was kept. In case of the forest development data, this

procedure resulted in forest development being not included in the model as a single term. Stats::anova() comparing the model with and without forest development as single term (next to soil depth and the two random terms) resulted in no significant difference (p = 0.46) between these models. Forest development did not have a significant effect on feeding activity either (car::Anova showed forest development p = 0.54 for the model containing forest development), therefore we kept the simpler model. In a later step of model selection, however, interactions of forest development with other factors were included in the model as they significantly improved the model. The significance level of variables and interactions was estimated via Type III sums-of-squares Wald Chi-square test using the Anova() function in the "car" package. The final models were as follows:

for hydrological regime:

```
"Feeding activity" ~ Soil_depth + Hydrology + Texture + PC2 + PC3 + Hydrology : Soil_depth + Hydrology : Texture + Hydrology : Season + Soil_depth : Season + Texture : Season + (1|SiteID) + (1|Season);
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and for forest development:

```
"Feeding activity" ~ Soil_depth + Soil_depth : Forest_development + Soil_depth : PC2 + Soil_depth : Season + Soil_depth : Texture + Forest_development : Texture + (1|SiteID) + (1|Season).
```

Pairwise comparisons of factors were calculated using the emmeans() function in the "emmeans" package (Lenth 2022). Since soil depth was treated as a numerical instead of factorial variable, pairwise comparisons including soil depth were performed using the lstrends() function in the "emmeans" package.

Furthermore, the average depth of faunal activity was evaluated using the Depth index (DI) (Gongalsky et al. 2004), calculated as:

$$DI = ((\sum n_i d_i) / N),$$

with DI being the mean depth of feeding activity, n_i the number of all pierced bait holes at depth i, d_i the respective soil depth [in mm] and N being the total number of pierced bait holes over all soil depths and test strips of one plot. Separate linear mixed effects models (lmer()) for the hydrological regime and forest development data set were built with DI as the response variable using the methods described above.

The final models for effects on the Depth Index were as follows:

for hydrological regime:

```
Depth Index ~ Hydrology + Hydrology : Season + (1|SiteID) + (1|Season);
```

and for forest development:

Depth Index ~ Forest development + Texture + Forest_development : Texture + Forest_development : Season + (1|SiteID) + (1|Season).

All statistical analyses were performed using R statistical software, version 4.1.2 (2021-11-01, "Bird Hippie", R Core Team (2021), with R studio version 1.4.1717. Graphical visualizations were created using the "ggplot2" package (Wickham 2016).

Results

Ecological site gradients

The hydrological regime was related to a number of abiotic site parameters (Table 1; Suppl. material 1: table S1A). Briefly, average number of flooding days per year was 5.2 in the active floodplain, but 0 in the seepage water zone and tributary floodplains. Leaf area index, i.e. density of the forest, increased with distance to the main river, while pH and average sand content of soil decreased. Within the active floodplain, the four forest development stages were rather similar in their environmental factors. However, they differed in leaf area index with old dense and young forests having higher canopy cover (high leaf area index) than old sparse forests (Table 2; Suppl. material 1: table S1B). Comprehensive information on ecological parameters of single sampling sites is given in Suppl. material 1.

Feeding activities in different hydrological regimes

The hydrological regime did not directly affect overall feeding activities (p = 0.11), and soil texture, soil parameters (represented by PC2) and vegetation parameters (represented by PC3) alone also did not significantly affect feeding activity. In contrast, soil fauna feeding activity decreased with soil depth in all hydrological regimes (p < 0.001, Table 3), and general feeding activity rates were higher in autumn than in spring (only included as random term in the model). Feeding activities increased from spring to autumn to a higher extent at shallow than at

Table 1. Site parameters in sampling sites according to hydrological regime (sampled 2020); Values represent means ± standard deviation.

Hydrology (2020) Soil temperature [°C]	Active floodplain spring autumn		Seepage v	vater zone	Tributary	
			spring autumn		spring autumn	
	9.57 ± 0.24	15.0 ± 0.71	10.59 ± 1.06	14.37 ± 0.28	10.39 ± 1.21	13.90 ± 0.25
Soil moisture [%H ₂ O]	15.24 ± 3.99	9.95 ± 1.97	18.61 ± 3.93	17.77 ± 5.6	25.23 ± 3.98	17.79 ± 4.18
pH (at 10 cm soil depth)	4.93 ± 0.17		4.13 ± 0.62		3.82 ± 0.11	
Leaf area index	2.21 ± 0.46		3.22 ± 0.61		4.83 ± 0.24	

Table 2. Site parameters in sampling sites according to forest development (sampled 2019); Values represent means ± standard deviation.

Forest development stage (2019) Soil temperature [°C]	Grassland spring autumn		Young forest spring autumn		Old dense forest spring autumn		Old sparse forest spring autumn	
	Soil moisture [%H ₂ O]	24.69 ± 5.51	14.44 ± 2.26	20.36 ± 3.20	10.87 ± 5.02	24.89 ± 4.66	17.43 ± 5.59	20.33 ± 1.54
pH (at 10 cm soil depth)	5.05 ± 0.25		5.15 ± 0.26		5.15 ± 0.46		5.24 ± 0.19	
Leaf area index	0		2.74 ± 0.33		2.97 ± 0.43		2.00 ± 0.30	

larger soil depths, resulting in a stronger decrease in feeding activity with soil depth in autumn (soil depth x season: p < 0.0001, Fig. 2). However, the decrease in feeding activity with soil depth was also affected by hydrology, as reflected in a stronger depth decrease in tributary sites than in the active floodplain or seepage water zone (interaction soil depth x hydrology: p < 0.0001; pairwise comparisons p = 0.0001 for tributary vs. seepage water zone, p = 0.0004 for tributary vs. active floodplain, but p = 0.88 for active vs. seepage water zone). Activity rates varied according to the season in sites of different soil textures, with sites on sandy soils showing higher activities than those on loamy soils in spring, but not in autumn (interaction soil texture x season: p < 0.0001). The interaction of soil texture and hydrology was not included in the model because all tributary sites were of loamy soil, unbalancing the experimental design in this respect. For comparison of soil fauna feeding activity of single sites see Suppl. material 4.

The Depth Index, i.e. average soil depth of soil fauna feeding activity, was significantly affected by hydrological regime (p = 0.001) with maximum average activity being at larger depths in the seepage water zone compared to tributary sites. In active floodplain sites, the Depth Index was similar to that of the seep-

Table 3. Analysis of Deviance, results for hydrological regime dataset (**A**) and forest developmental stage dataset (**B**), separated into general feeding activity and Depth.

	Chisq	Df	Pr (>Chisq)	
A) Hydrological regime				
Feeding activity				
Soil depth	414.8	1	<0.0001 ***	
Hydrology	4.4	2	0.11	
Soil texture	0.2	1	0.65	
PC2 (soil parameters)	0.02	1	0.86	
PC3 (vegetation & soil pH)	3.4	1	0.07	
Soil depth: Hydrology	22.7	2	< 0.0001 ***	
Hydrology: Season	47.3	3	< 0.0001 ***	
Soil depth: Season	28.5	1	< 0.0001 ***	
Texture: Season	19.7	1	< 0.0001 ***	
Depth index				
Hydrology	13.6	2	0.0011 **	
Hydrology: Season	5.8	3	0.12	
B) Forest development				
Feeding activity				
Soil depth	39.9	1	<0.0001 ***	
Soil depth: Forest development	10.6	3	0.0130 *	
Soil depth: PC2 (Soil parameters)	2.3	1	0.1304	
Soil depth: Season	28.1	1	<0.0001 ***	
Soil depth: Soil texture	13.8	1	0.0002	
Forest development: Season	36.1	4	<0.0001 ***	
Forest development: Soil texture	39.4	3	<0.0001 ***	
Depth index				
Forest development	23.3	3	<0.0001 ***	
Soil texture	6.4	1	0.0113 *	
Forest development: Soil texture	10.8	3	0.0126 *	
Forest development: Season	18.4	4	0.0010 *	

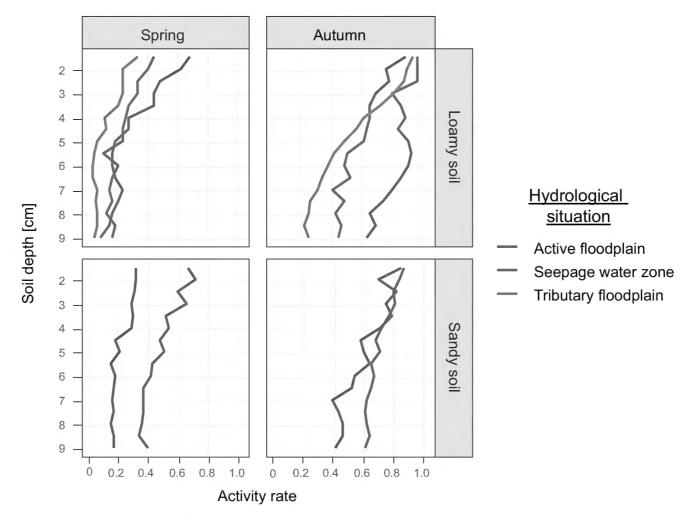


Figure 2. Feeding activities in different hydrological situations separated by season and soil texture.

age water zone, but the difference to tributary sites was not significant (Fig. 3, Table 3A). Neither season nor soil texture nor other environmental variables (PCs) modified this effect.

Feeding activities in the forest development stages

Within the active floodplain, forest developmental stage, soil texture, PC2 and PC3 were not included in the final model as single factors (only as interaction terms) due to higher AIC values in models containing these variables. Therefore, the effect of these variables individually was considered to be non-significant. As in the hydrological gradient, soil fauna feeding activity decreased with soil depth in all forest development stages (p < 0.0001, Table 3B). Further, the decrease in feeding activity with soil depth was more pronounced during spring than autumn (soil depth x season interaction p < 0.0001), and in sandy compared to loamy soils (p < 0.0001 for spring vs. autumn and p = 0.0004 for loamy vs. sandy, Fig. 4). General feeding activity rates were higher in autumn than in spring (only included as random term in the model), but significant differences between forest developmental stages were only observed in spring (forest developments x season interaction p < 0.0001) with lower feeding activity rates in grasslands than in old dense and young forests in spring.

Further, in grasslands on sandy soil, feeding activity was lower than in dense old and young forests on sandy soil (p = 0.01), while on loamy soil the differences between forest developmental stages were not significant. In addition, in grasslands on sandy soils feeding activity rates decreased stronger with soil depth than in young forests (interaction soil depth x forest development: p = 0.01; comparison young vs. grassland across depth with p = 0.029; all other comparisons p > 0.05). For soil fauna feeding activity in the individual sampling sites, see Suppl. material 5.

Forest development significantly affected Depth Index, i.e. maximum average soil-fauna activity occurred at a larger average soil depth in young forests than

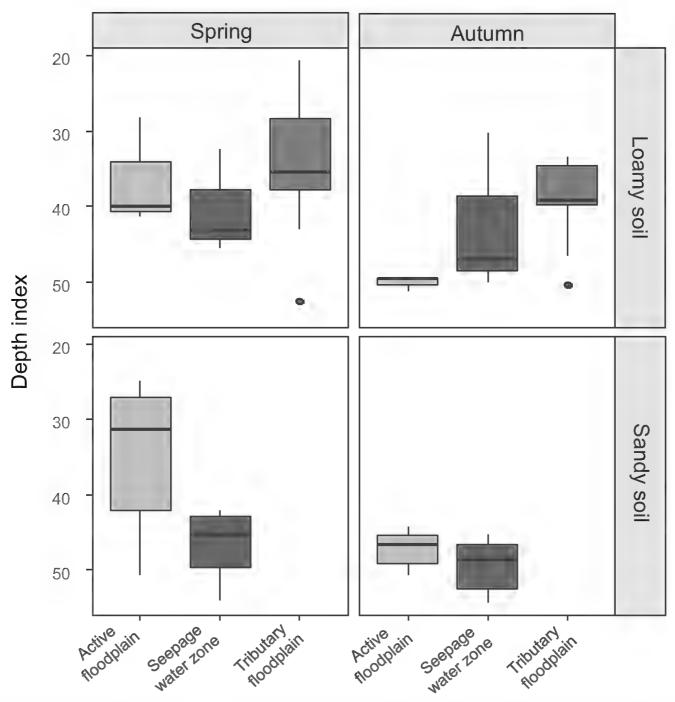


Figure 3. Depth index after Gongalsky et al. (2004) for hydrological situations according to season and soil texture; all investigated tributary sites were of loamy soil texture, therefore no value for sandy tributaries. Note the inverted y-axis for better imaginability of feeding depth.

in grasslands (p = 0.04), while average soil depth of the activities in old dense and old sparse forests were between these values (Table 3B, Fig. 5). Soil texture affected Depth Index, resulting in maximum average soil-fauna feeding activity at greater soil depths in loamy compared to sandy soils. The effect of forest development on the Depth Index was modified by season and soil type. While in spring the mean Depth Index was lower in grasslands than old dense and young forests (but not old sparse forests), no difference was found in autumn. While no significant difference was found between forest development stages on loamy soils, the mean Depth Index was significantly lower in grasslands compared to young, old dense and old sparse forests on sandy soils (p = 0.03).

Discussion

General effects

The feeding-activities observed here for the most part conform to other studies, confirming the suitability of the Bait-Lamina test in the current study, although adjustment of exposure time would have been necessary during autumn 2019 (see below). Soil-fauna feeding activity significantly decreased with soil depth and increased from spring to autumn in both gradients, similar to findings of

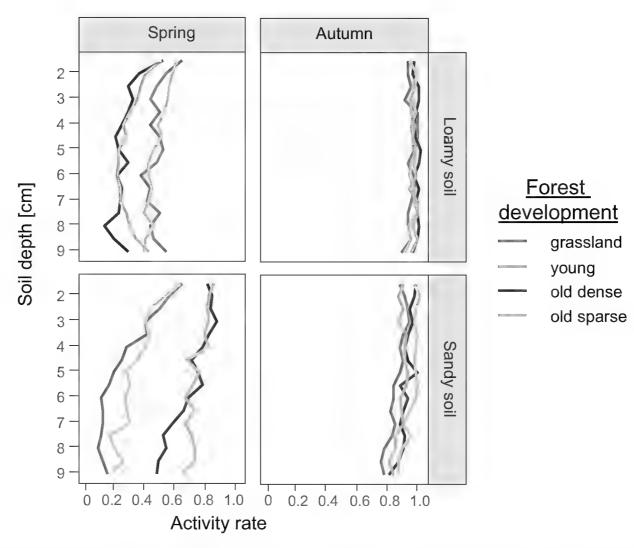


Figure 4. Feeding activities in different forest development stages separated by season and soil texture.

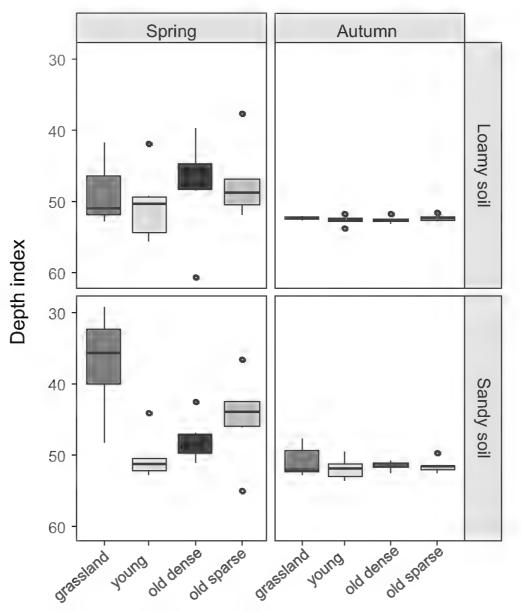


Figure 5. Depth index after Gongalsky et al. (2004) for different forest development stages according to season and soil texture; Note the inverted y-axis for better visualisation of feeding depth.

earlier studies (e.g., Geissen et al. 2007; Rozen et al. 2010). This was expected, since decreasing soil fauna density with soil depth results in lower measured activity in deeper soil (Birkhofer et al. 2011). Soil fauna feeding activity has also

been observed to increase with temperature (Gongalsky et al. 2004; Rozen et al. 2010; Simpson et al. 2012; Klimek et al. 2015), suggesting potential higher activity in early autumn and at sun-exposed soil surfaces.

Nonetheless, in the forest development gradient, feeding activities in autumn were close to 100% in all forest and grassland sites (see Fig. 4), indicating that the ISO-recommended exposition time of 4 weeks during autumn was too long to reveal differences in feeding activity. In the hydrological gradient, on the other hand, four weeks exposition was appropriate in early autumn, since the measured activities remained below 90%, suggesting that the slightly lower temperatures in this sampling year reduced overall activity (Suppl. material 1 for comparison of soil temperatures). Interestingly, soil fauna density and species richness did not significantly influence observed feeding activities, confirming previous reports that such community parameters cannot serve as proxies for feeding activity (see Introduction).

Hydrological regime

Besides frequency of inundation, the hydrological regime was strongly associated with further environmental factors, separating (1) forests of low leaf-area index, i.e. sparse canopy cover, on soils of neutral pH in the active floodplain from (2) very dense forests on strongly acidic soils in the tributary floodplains and (3) forests in the seepage water zone with intermediate canopy cover and soil pH. The factor "hydrological regime" therefore represented a number of environmental factors and was expected to significantly affect soil fauna feeding activity as well. However, a general overall effect was not confirmed and no significant differences in average feeding activity were found between hydrological regimes. Soil fauna feeding activity was characterized by high variability between different sites of the same hydrological regime, indicating that further factors besides general inundation type (and associated environmental conditions) affected soil fauna activities in our study. In addition, since no severe flooding event in the Elbe floodplain had occurred for 6 years prior to our study, we suggest that no effect of direct inundation persisted in the active floodplain sites, which would be in line with findings by Russell (2008) showing changes in the soil fauna community structure persisting only for several months after an inundation event.

Although overall feeding activity of soil fauna did not differ between hydrological regimes, the decrease in activity with soil depth was significantly stronger in tributary sites than sites of the seepage water zone or active floodplain, in particular in autumn. This indicates that soil-fauna activity was prevented deeper in the soils of the tributary sites, but not in the other hydrological regimes (cf. Rozen et al. 2010). High ground-water levels and resulting anoxia in deeper soil layers are a plausible explanation for restricted faunal vertical movement – and thus their low activity – in these sites. It is remarkable that anoxia by high ground water level had occurred mainly in the floodplain of tributaries, but not in the active floodplain, which contradicts our expectations. We therefore assume that the stagnant water in sites behind the dyke is much more negatively affecting soil fauna feeding activities than the running (flood) water in the active floodplain. In addition, anoxic conditions reduce mobility of mobile nutrients and increase toxicity of aluminium (Geissen and Brümmer 1999; Geissen et al. 2007), thereby potentially reducing soil fauna density and activity. In our study,

hydromorphic soil characteristics indicated that anoxic conditions – due to a high groundwater table – occurred at rather shallow soil depths in sites of tributaries as well as seepage water zone (at 35 and 40 cm, respectively, compared to > 100 cm in the active floodplain soils; Vásconez Navas et al. 2023). Surprisingly, a reduction of soil fauna activity in deeper soil only occurred in tributary floodplain sites but not the seepage water zone, indicating at least temporal differences in soil water saturation despite occasional anoxic conditions in both hydrological regimes. All investigated tributaries are regulated for agricultural purposes (pers. comm. A. Gröngröft). We therefore assumed a complete independence of water levels between the Elbe River and its tributaries with anoxic conditions by high groundwater table probably persisting throughout the summer in the tributary floodplain soils. In sites of the active floodplain and seepage water zone, on the contrary, low water levels of the Elbe River in summer resulted in lower groundwater tables, allowing soil fauna to migrate deeper. Consequently, our results indicate a higher impact of recent hydrological effects (i.e. high groundwater tables) than of long-term inundation effects.

Soil texture alone apparently did not affect soil fauna feeding activity in our study, but the absence of sandy tributary sites hampered a comparison of soil-texture effects between all hydrological regimes. In the seepage water zone higher soil fauna activities were observed in sandy compared to loamy soils. This indicates better abiotic conditions for the soil fauna in sandy soils in the seepage water zone with probably better aeration and warming of the soil surface, in spite of the lower water holding capacity and therefore lower mean soil moisture of sandy soils (Bardgett 2005). However, this effect was restricted to the seepage water zone, since activity rates were similar in sandy and loamy soils in sites of the active floodplain, suggesting that in the active floodplain different parameters might affect soil fauna activity. Furthermore, when comparing activity rates in sites of the seepage water zone to those of the active floodplain, the positive effect of sandy soil becomes only visible during spring. During early autumn, on the contrary, activity rates were similar in the active floodplain and sandy seepage water zone sites, but there was a negative effect of loamy soil in sites of the seepage water zone. Our sampling campaigns took place in years of extreme drought in Europe, leading to very dry soils in the investigated sites. In particular, in the seepage water zone sites with loamy soil, very low soil water content was observed, and in addition loamy soils became very dry and hard during summer. Therefore, the positive impact of increasing temperature on soil fauna during early autumn probably was limited by a decrease in soil moisture, as soil fauna is known to avoid dry conditions by retreating to deeper soil horizons or inactive stages (Gongalsky et al. 2008; Thakur et al. 2018). In conclusion, abiotic soil conditions seemed to affect soil fauna activity in particular in the seepage water zone, indicating a limitation of soil fauna feeding activity by soil temperature in spring, but by soil moisture in autumn.

Forest development gradient

Since in the present study all sites of the forest development gradient were located in the active floodplain, they represented a habitat diversification in this specific hydrological regime. However, environmental and vegetation parameters varied mainly between grassland sites on the one hand and forest sites on the other hand.

This was expected to be reflected in the observed soil fauna activity. Earlier studies did not find differences in soil fauna feeding activity between central European coniferous and deciduous forests of different ages (Geissen et al. 2007; Klimek et al. 2015), but between grasslands and forest sites in general (Hamel et al. 2007). In our study, forest development also did not directly affect soil fauna feeding activity and we found neither a significant difference in overall feeding activity between grasslands and forests, nor between forests of different age or density. Instead, we found that forest development modulated the effect of soil depth, i.e., a steeper decrease of activity with soil depth was observed in grasslands than in young forests. Differences in soil fauna feeding activity between grasslands and young forests only occurred at > 2.5 cm soil depth, and DI values suggested that in young forests mean soil fauna activity occurred at a greater soil depth as compared to grasslands. We therefore conclude that in grasslands soil fauna activity was restricted to the soil surface by environmental factors, probably by rapidly declining soil moisture with soil depth due to intense water uptake of grass roots (pers. obs.). On the other hand, in forests (in particular young forest plantations) soil moisture was likely appropriate for soil fauna feeding activity at least to 9 cm depth.

Soil texture also did not affect soil fauna feeding activity directly in this gradient and, due to very high general feeding activities in autumn, significant differences between soil fauna activities in different forests or grassland sites were only observed in the spring sampling. In old dense and young forests, feeding activity was higher in sandy compared to loamy sites. The most probable explanation for this result is the lower exposition of the soil to sunlight in denser forests. Old dense and young forests exhibited lower soil temperatures (mean difference of ~0.5 °C) and higher moisture (mean difference of ~5%) in loamy sites compared to sandy sites, while in open habitats (sparse forests and grasslands) these differences were less pronounced. This indicates that in shady habitats, sandy soils provided better habitat conditions for soil fauna than loamy soils due to faster warming (before development of canopy cover during spring) and reduced potential anoxia, while the close canopy cover prevented desiccation of the soil surface (Thakur et al. 2018). On the contrary, in grasslands and old sparse forests the effect of soil texture was restricted to deeper soil with a marked decrease in soil fauna feeding activity in sandy but not in loamy soils. This was surprising since the higher porosity of sandy soil, resulting in higher oxygen content in deeper soil, was expected to increase soil fauna activity in deeper soil (cf. Hassink et al. 1993). Soil moisture and temperature were similar in loamy and sandy soils in open habitats, indicating that the causes of similar feeding activity rates at the soil surface, but differences in deeper soil, might not lie in exposure to sunlight. Instead, differences most probably were related to site-specific effects e.g. pH or groundwater level, that resulted in strong differences in feeding activity rates between sites of individual forest developmental stages (Suppl. material 5).

Conclusion

In the present study we found soil fauna feeding activity to be much more affected by current than average (concluded from hydromorphic soil characters) groundwater levels. This indicates that soil fauna activity rates fluctuate with time and react to variable conditions in soil moisture and/or soil temperature within short time. Further, varying conditions within the floodplain, depending

mainly on vegetation cover, shape environmental conditions. The soil fauna feeding activity in floodplain hardwood forests seems to be limited by soil temperature during spring, but by soil moisture during summer with hydrological regime and forest density affecting soil fauna feeding activity only indirectly via influencing vegetation cover and therefore soil temperature and moisture.

Our study made a first attempt to assess the effects of environmental conditions on soil fauna activity in hardwood floodplain forest-restoration sites, showing that a combination of various site conditions is highly influential for faunal activity. Future, more specific studies are needed to understand the mechanisms of these effects and investigate the contribution of soil fauna to ecosystem services such as decomposition and carbon cycling in floodplain habitats, as well as to understand the detailed relationships between soil fauna density and diversity, environmental conditions, and feeding activity. Our results indicate dense forests on sandy soil, located in the active floodplain or seepage water zone, having the highest potential value for the ecosystem service of organic-matter decomposition. From a nature conservational perspective, these could be preferred sites for floodplain forest restoration in current grassland.

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Additional information

Conflict of interest

The authors have declared that no competing interests exist.

Ethical statement

No ethical statement was reported.

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Author contributions

Conceptualization: DJR. Data curation: NS. Formal analysis: NS. Funding acquisition: DJR. Investigation: NS. Methodology: NS, DJR. Project administration: NS, DJR. Resources: DJR. Supervision: DJR. Validation: NS. Visualization: NS. Writing – original draft: NS. Writing – review and editing: NS, DJR.

Author ORCIDs

Nicole Scheunemann https://orcid.org/0000-0003-1845-6236 David J. Russell https://orcid.org/0000-0002-0129-0375

Data availability

All of the data that support the findings of this study are available in the main text or Supplementary Information.

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Supplementary material 1

Soil parameters of all study sites

Author: Nicole Scheunemann

Data type: pdf

Explanation note: Sampling sites in different hydrological regimes differed in the number of flooding days per year: sites in the active floodplain had been flooded at an average of 5.2 days per year between 1990 and 2016, while sites in the seepage water zone and floodplain of tributaries had not been flooded by the Elbe. The exception was one site in the seepage water zone that had been flooded at a single dyke breach during a historic flood event in 2002. All sites of the seepage water zone had been exposed to rising ground water tables depending on the Elbe water gauge, while groundwater tables in the floodplain of tributaries were in dependent of the Elbe flood regime (but were exposed to that of the respective small tributary). Sites differed mainly by leaf area index, pH and sand content, i.e. forests in the active floodplain forests were sparse and soils were of neutral pH and medium sand content, while forests in the seepage water zone were more dense and soils had more acidic pH and high sand content. Forest in the floodplain of tributaries were most dense and showed low pH and comparably low sand content. Soil moisture was lowest in the active floodplain, but the decrease of soil moisture from spring to autumn was most pronounced in tributary sites. Within the active floodplain, sites sampled in the forest development gradient mainly differed by forest age (156.8 years in old dense forests, 130.8 years in old sparse forest, 22.3 years in young forests and 0 in grasslands). In most aspects, the three developmental stages with growing trees were rather similar in their environmental factors, e.g. number of plant species and cover of vegetation layers, litter cover, pH, sand content of soil, mean soil temperature, etc. However, they differed in leaf area index with old dense and young forests being more shadowed (high leaf area index) than old sparse forests. Surprisingly, grasslands had an average higher number of flooding days per year (49 in grasslands as compared to a maximum of 28 in forests). As expected, vegetation parameters differed between grasslands and forests, but most other environmental variables were similar to those in forests.

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Supplementary material 2

PCA plot of sites and environmental parameters of the hydrology gradient

Author: Nicole Scheunemann

Data type: pdf

Explanation note: PCA of ecological site parameters for hydrological regimes sampling 2020; Gray triangles indicate factor coordinates of closest environmental variable. Site IDs in bold with frame with color indicating hydrological situation: turquoise = active floodplain, purple = seepage water zone, red = tributary floodplain.

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Supplementary material 3

PCA of ecological site parameters of all sites in the forest development gradient

Author: Nicole Scheunemann

Data type: pdf

Explanation note: PCA of ecological site parameters of forest development sampling 2019; gray triangles indicate factor coordinates of closest environmental variable. Site IDs in bold with frame, with color indicating forest development stage: dark blue = old dense forest, light blue = old sparse forest, gray = young forest plantation, orange = grassland.

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Supplementary material 4

Soil fauna feeding activity of single sites in hydrology gradient

Author: Nicole Scheunemann

Data type: pdf

Explanation note: Soil fauna feeding activity of single sites (means of three replicates) of hydrological situations in spring (above) and autumn (below) 2020; line types represent site ID, legend is valid for spring as well as autumn sampling.

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Supplementary material 5

Soil fauna feeding activity of single sites in forest development gradient

Author: Nicole Scheunemann

Data type: pdf

Explanation note: Soil fauna feeding activity of single sites (means of three replicates) of forest development stages in spring (above) and autumn (below) 2019; Line types represent site ID, legend is valid for spring as well as autumn sampling.

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